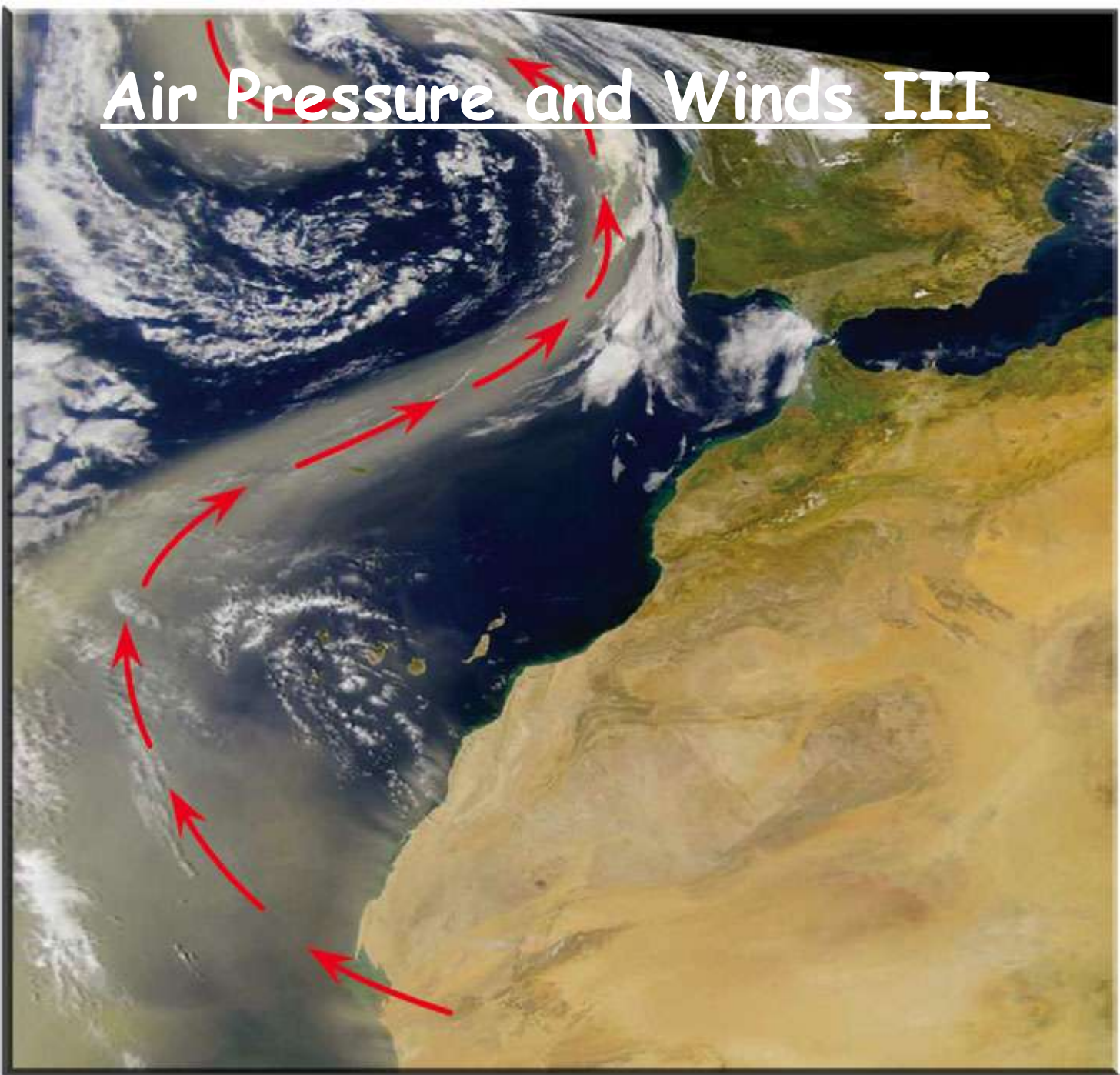


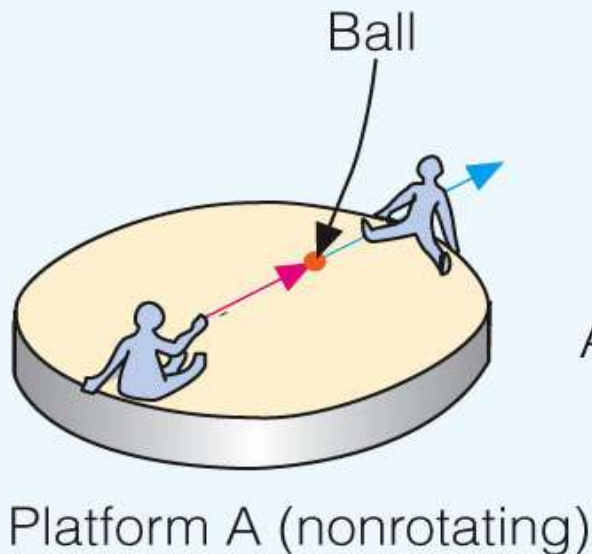
Air Pressure and Winds III



Coriolis Force (Effect)

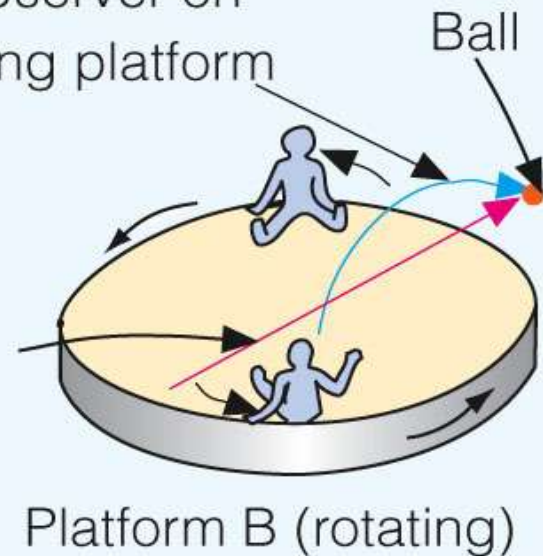


- It is an **apparent** force;
- Due to the **rotation** of the coordinate system (Earth);
- It makes a moving object **deflect** from a straight line even in the absence of any forces acting on it.



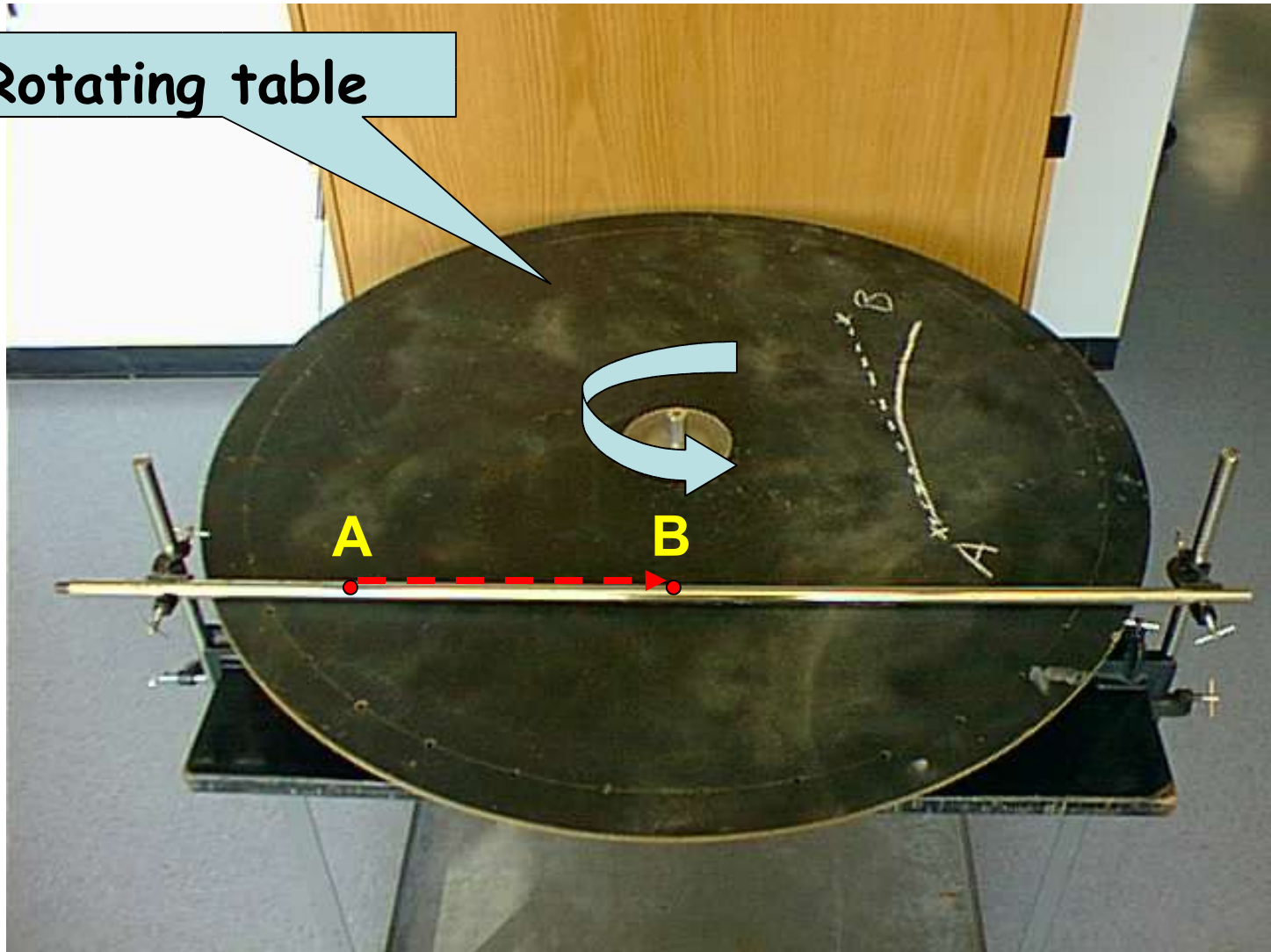
Actual path

Apparent path as seen by observer on rotating platform



Coriolis Force Demonstration

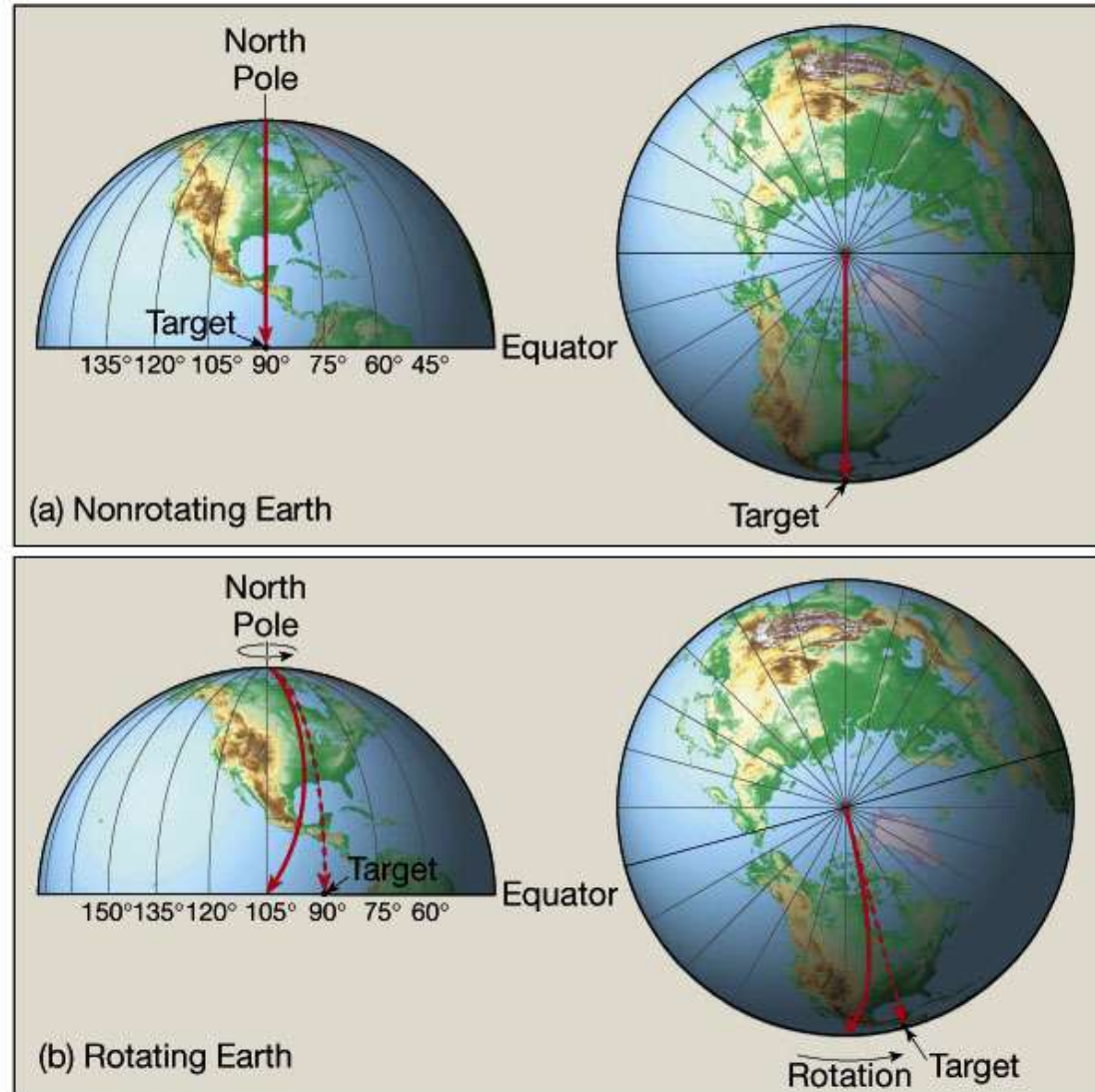
Rotating table



Dashed line - the trajectory of the chalk with respect to a non-rotating table.
Solid line - the trajectory of the chalk with respect to a rotating table.

The Magnitude of the Coriolis Force

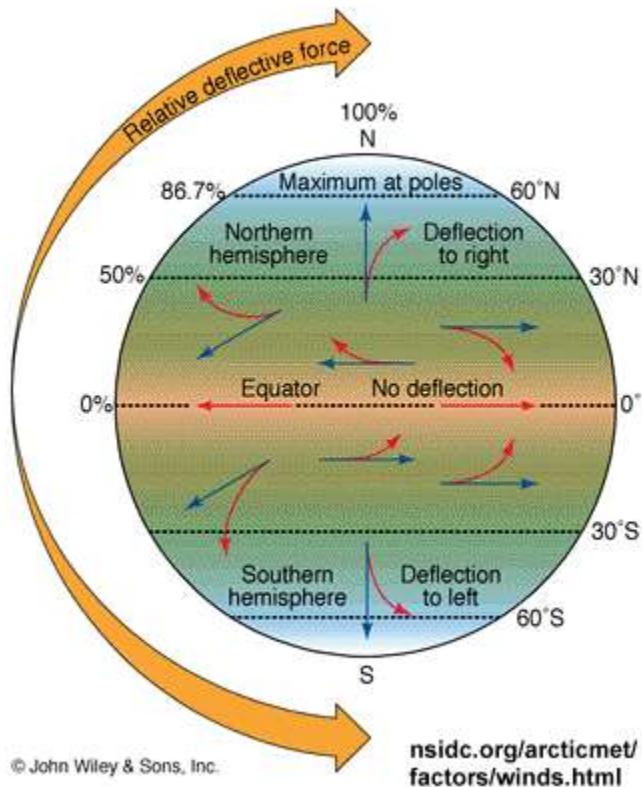
- The **rotation** of the Earth
 - ◆ The faster the planet rotates the bigger the force
- The **speed** of the object
 - ◆ Bigger $V \rightarrow$ bigger effect
- The **latitude**:
 - ◆ Min. at the equator
 - ◆ Max. at the poles



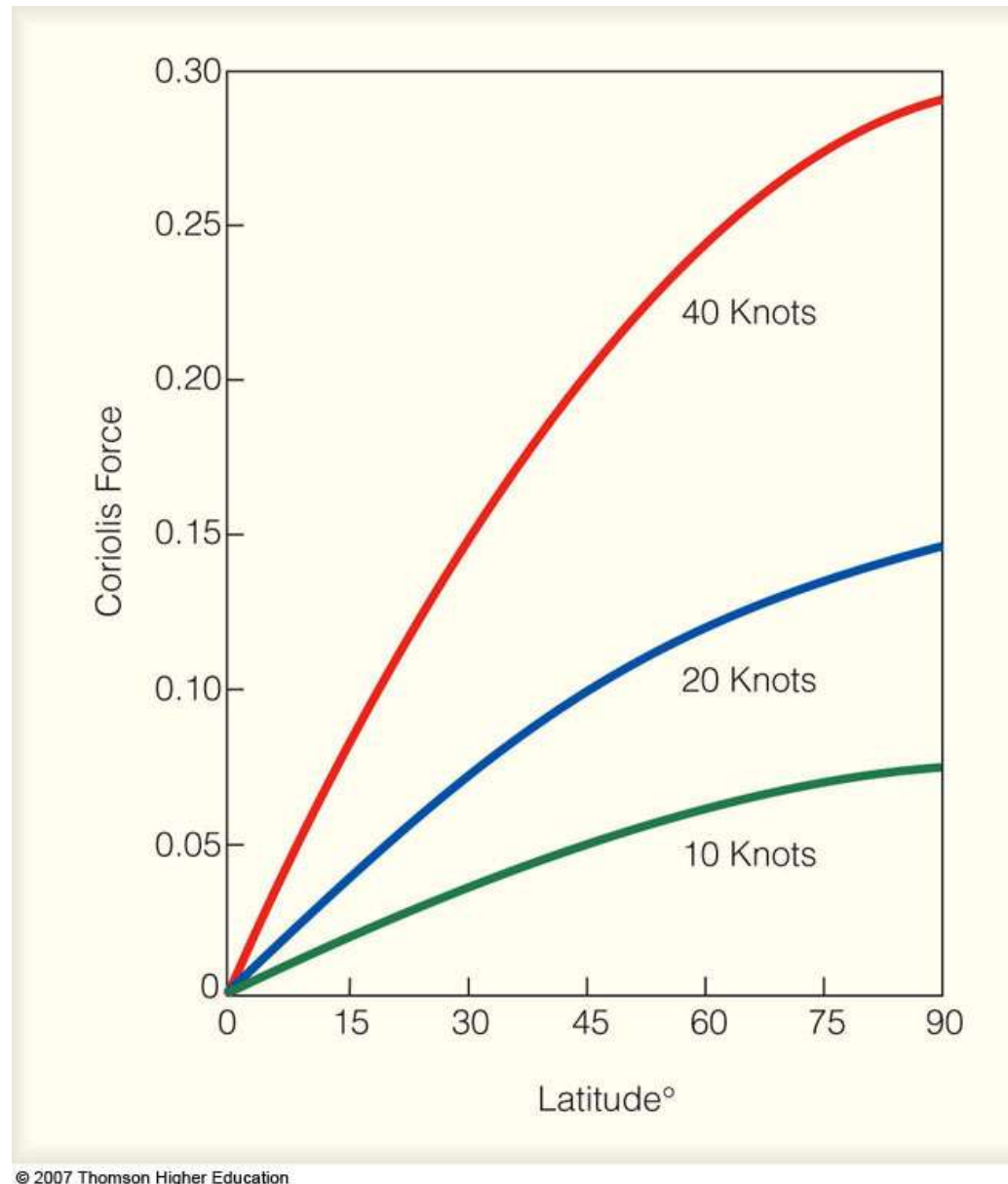
$$F_{co} = 2m\Omega V \sin \varphi$$

Coriolis force as a function of:

- The **speed** of the object
- The **latitude**:
 - ◆ Min. at the equator
 - ◆ Max. at the poles



$$F_{co} = 2m\Omega V \sin \phi$$

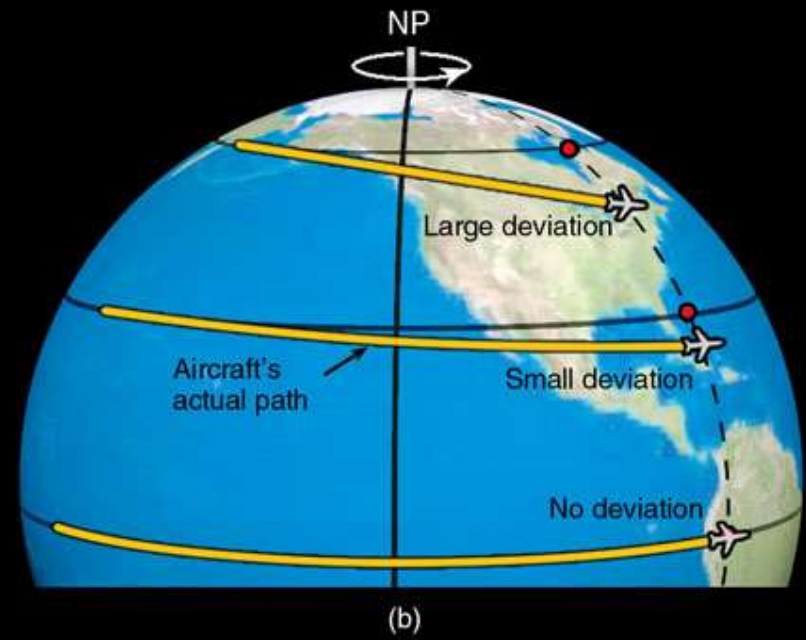
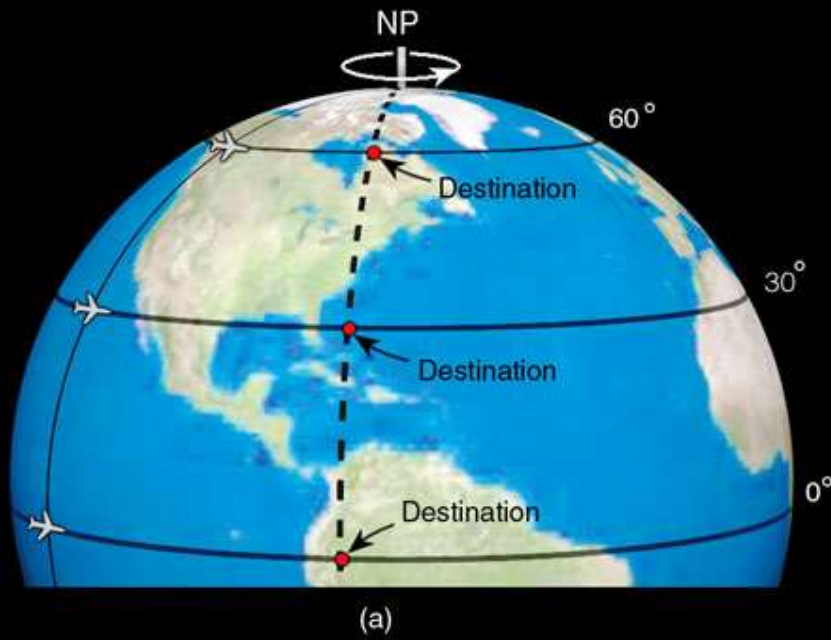


The Direction of the Coriolis Force

- In the **Northern hemisphere** the deflection is to the **right** of the direction of motion.
- In the **Southern hemisphere** the deflection is to the **left** of the direction of motion.
- The winds in the Northern hemisphere will be deflected to the right and in the Southern hemisphere they will be deflected to the left.
 - ◆ Hurricanes spin differently in the Northern and Southern hemisphere

The Coriolis Force and the Earth

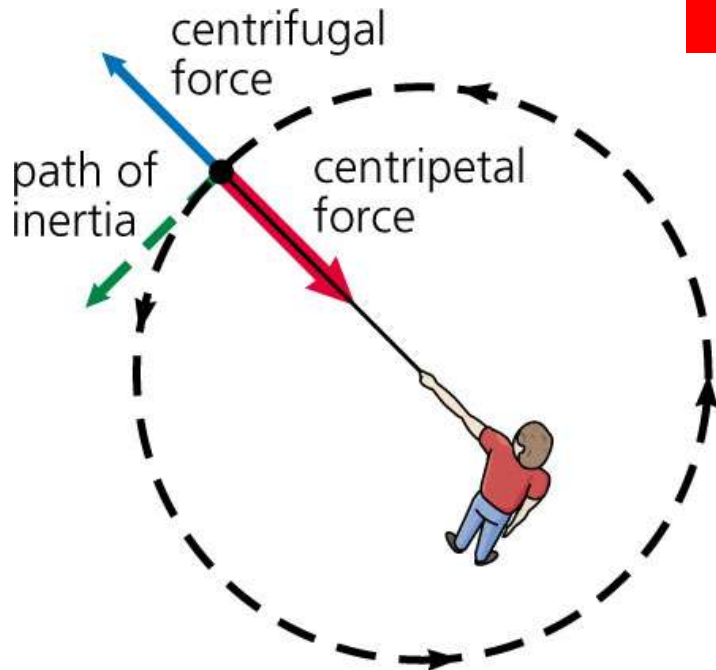
- The Coriolis effect is important when moving over **LARGE** distances (air plane travel), with large velocities, away from the equator.



Centripetal/Centrifugal Force

- Any motion in a **curved path** represents accelerated motion, and requires a **force** directed toward the center of curvature of the path. This force is called the **centripetal force** which means "center seeking" force.

$$F = ma = m \frac{v^2}{r}$$



Properties of the Centripetal Force (CF)

- The centripetal acceleration, and the centripetal force are **perpendicular to the direction of motion**.
- They only change the **direction of motion**.
- They do **NOT** change the magnitude of the velocity.
- The **CF** changes the direction of the wind but not the magnitude of the wind.



Recap: Forces in the Atmosphere

- Gravity force.

- ◆ Vertical force in a downward direction

$$G = mg$$

- Atmospheric drag force (friction).

- ◆ Acts against the motion
- ◆ Proportional to velocity squared

$$F_{drag} = \frac{1}{2} C_p \rho_{air} A v^2$$

- Pressure gradient force

- ◆ From high to low pressure regions
- ◆ Perpendicular to the **isobars**
- ◆ The bigger the pressure gradient (denser the isobars), the larger the pressure force

$$F_p \propto - \text{pressure gradient}$$

- Coriolis force: due to the Earth's rotation

- ◆ Deflection **to the right** in the Northern hemisphere
- ◆ Varies with latitude (absent at equator, max at the poles)
- ◆ Proportional to the velocity of the object (wind)

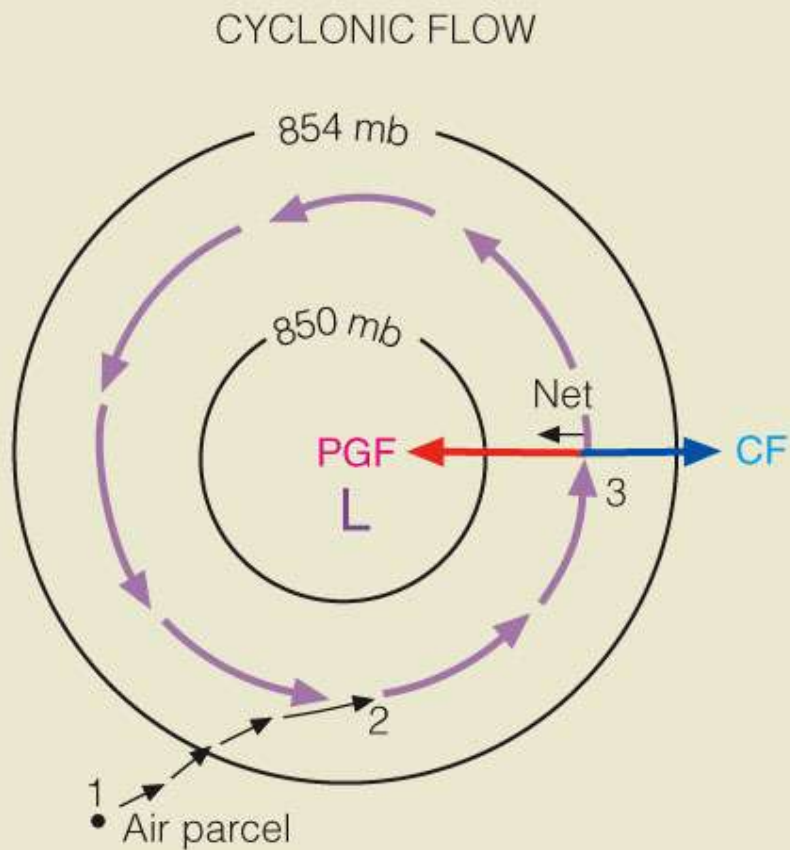
$$F_{co} = 2 m \Omega V \sin \varphi$$

- Centripetal force:

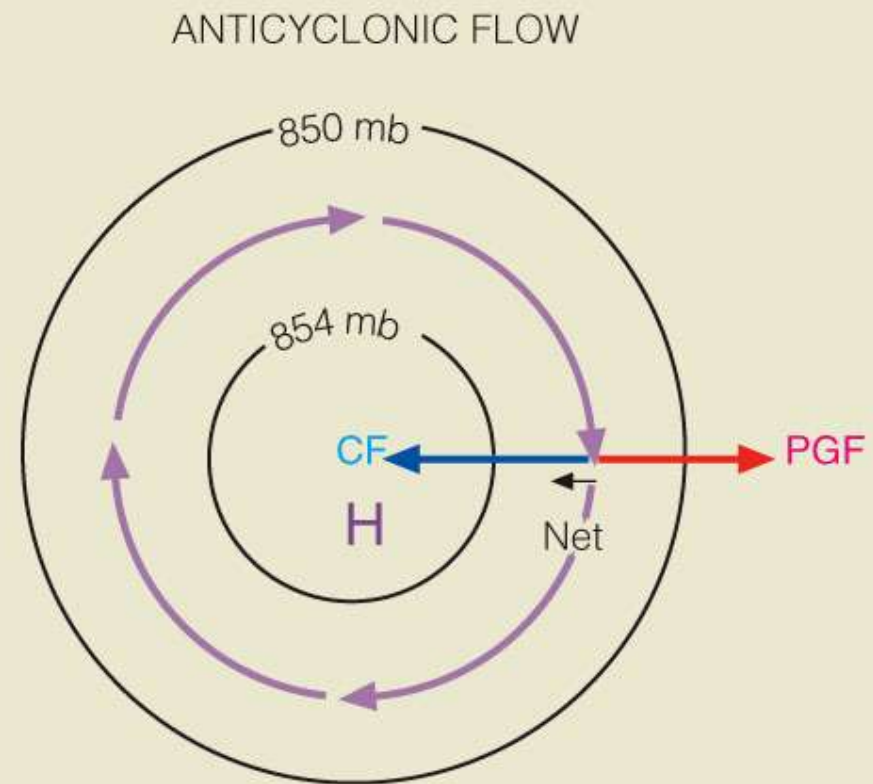
$$F \propto \frac{V^2}{r}$$

Origin of the Centripetal Force

- The centripetal force in the atmosphere is the net result of the **pressure gradient force** and the **Coriolis force**

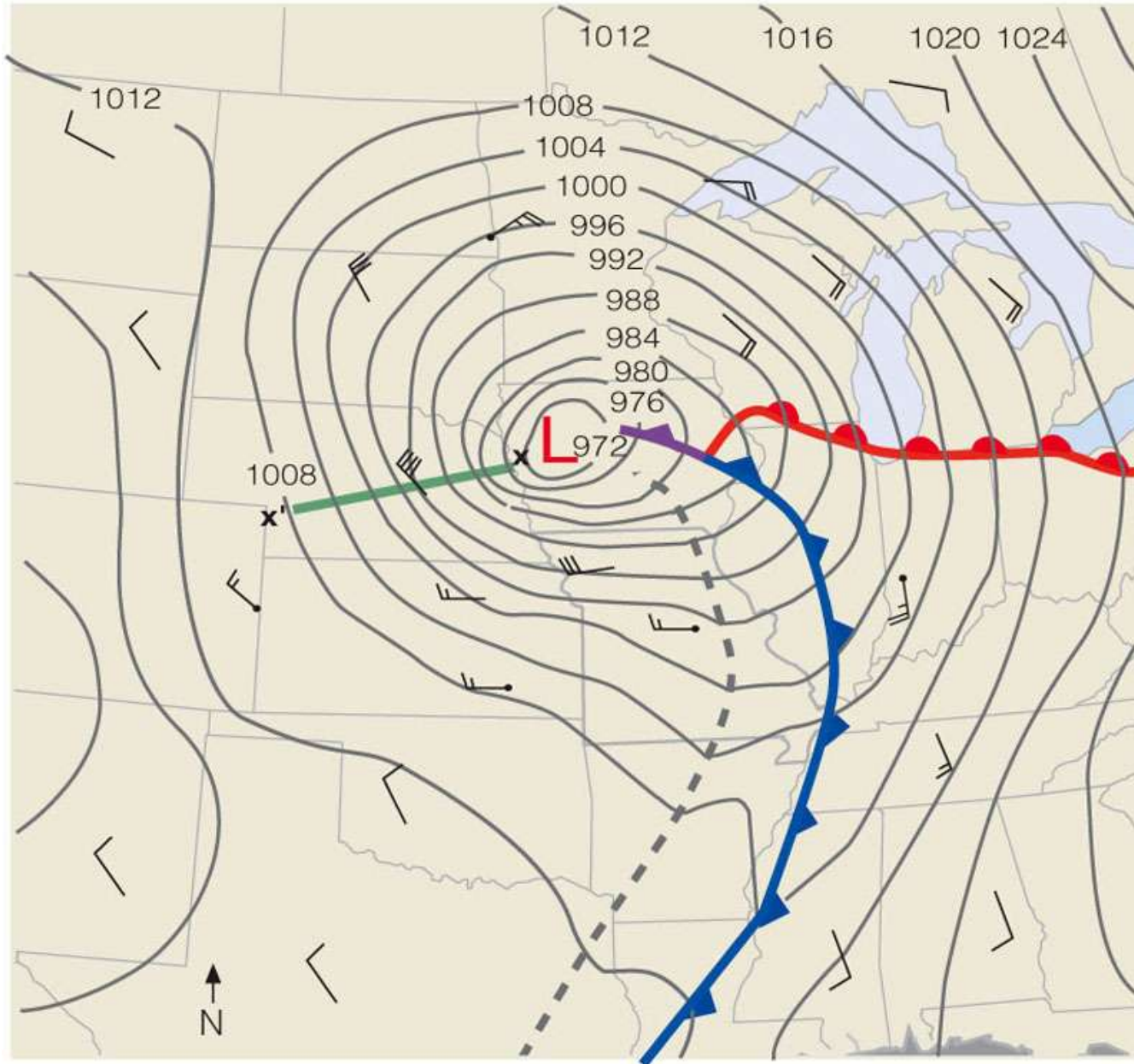


(a) Low pressure area (cyclone) aloft



(b) High pressure area (anticyclone) aloft

Wind and Pressure Map



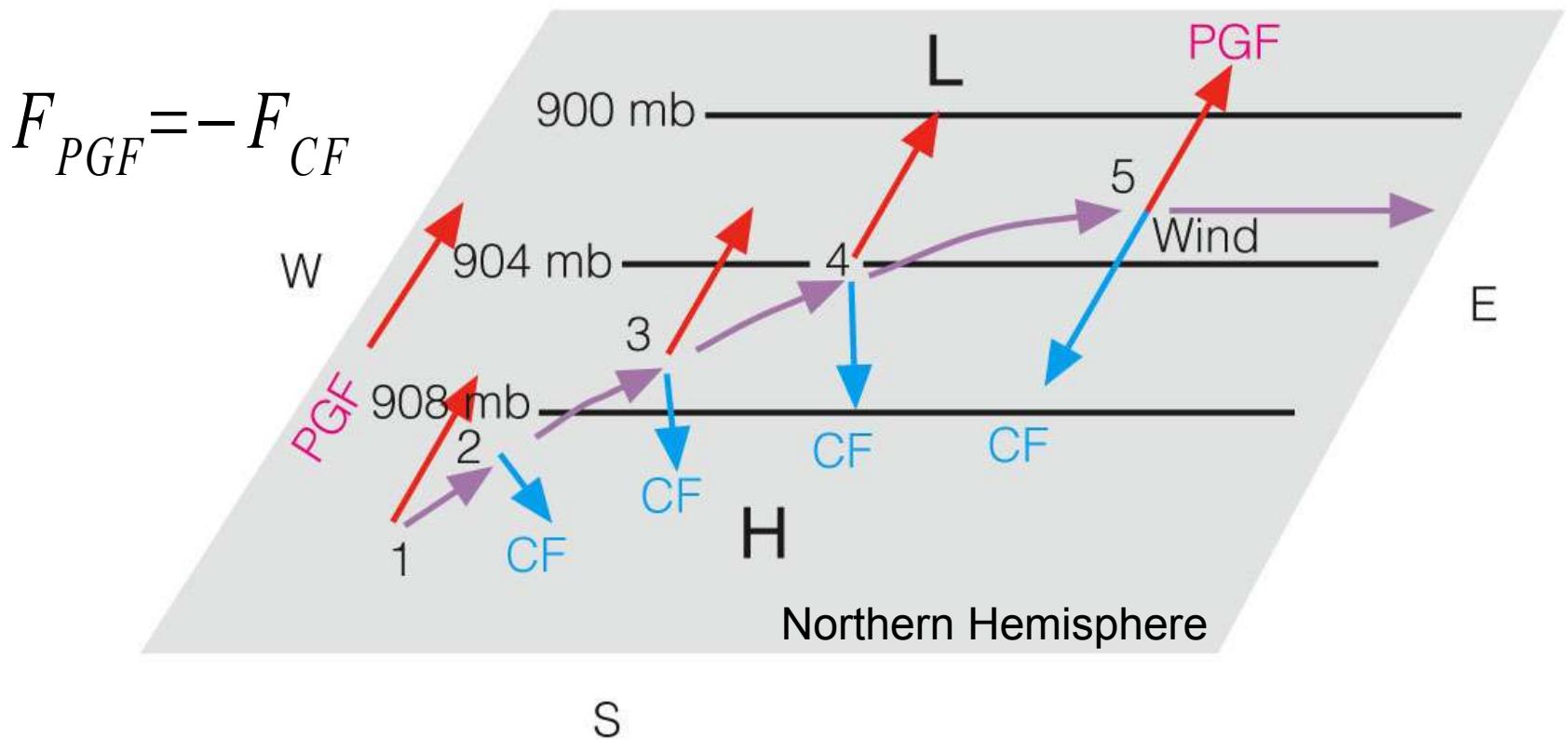
	Miles (statute) per hour	Knots
☉	Calm	Calm
—	1–2	1–2
↖	3–8	3–7
↖	9–14	8–12
↖	15–20	13–17
↖	21–25	18–22
↖	26–31	23–27
↖	32–37	28–32
↖	38–43	33–37
↖	44–49	38–42
↖	50–54	43–47
↖	55–60	48–52
↖	61–66	53–57
↖	67–71	58–62
↖	72–77	63–67
↖	78–83	68–72
↖	84–89	73–77
↖	119–123	103–107

Winds in the Atmosphere

- Geostrophic winds
 - ◆ Pressure gradient force = Coriolis force
- Gradient winds
 - ◆ Pressure gradient force not equal to Coriolis force
 - ◆ Cyclones: $PGF > CF$
 - ◆ Anticyclones: $PGF < CF$
- Surface winds
 - ◆ Affected by ground friction
- Vertical air motion

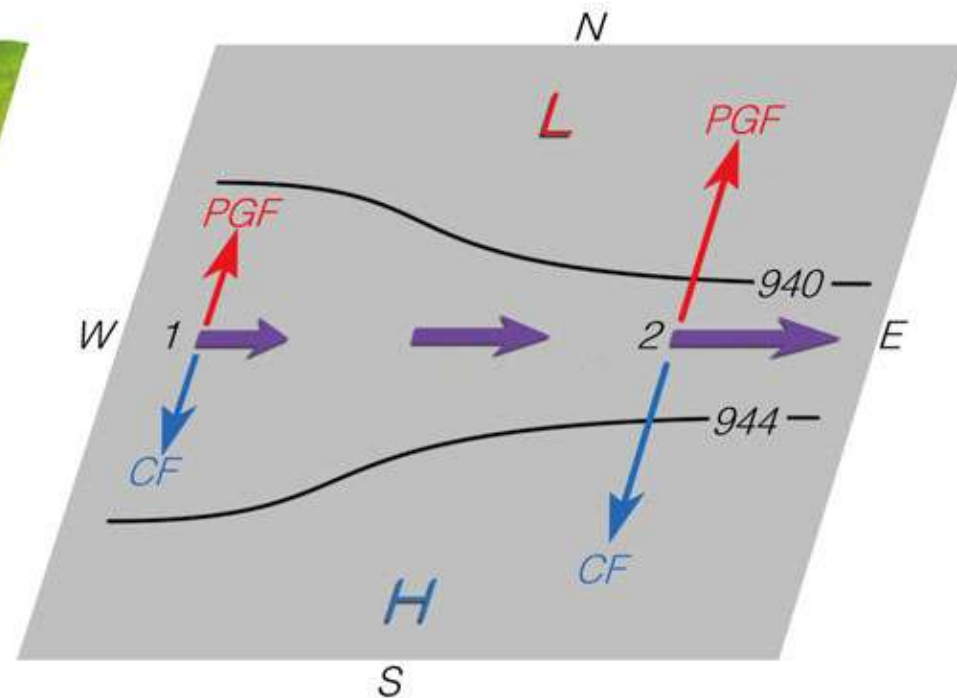
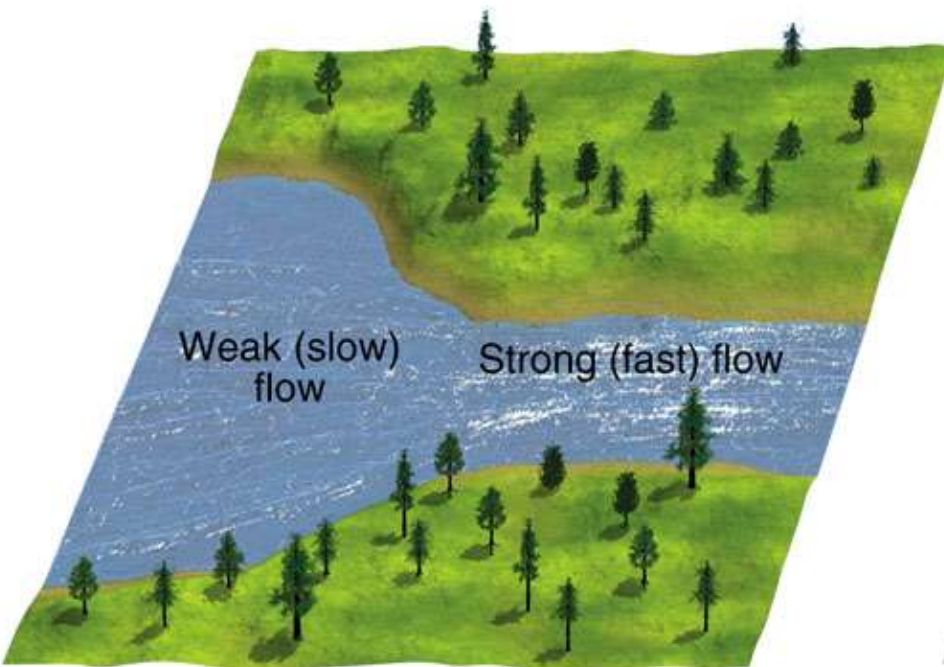
Geostrophic Winds: direction

- The **pressure gradient force** balances the **Coriolis force**.
- Typically occur at **higher altitudes** (>1 km).
- The winds are **parallel to the isobars**.
- In the **NH** the low pressure is to the **left** of the wind direction and in the **SH** the low pressure is to the **right**.



Geostrophic Winds: speed

- The wind speed is proportional to the density of the isobars - analogy to a water in a stream
 - ◆ Density of isobars increases \rightarrow PGF increases
 - ◆ Wind speed increases \rightarrow CF increases as well



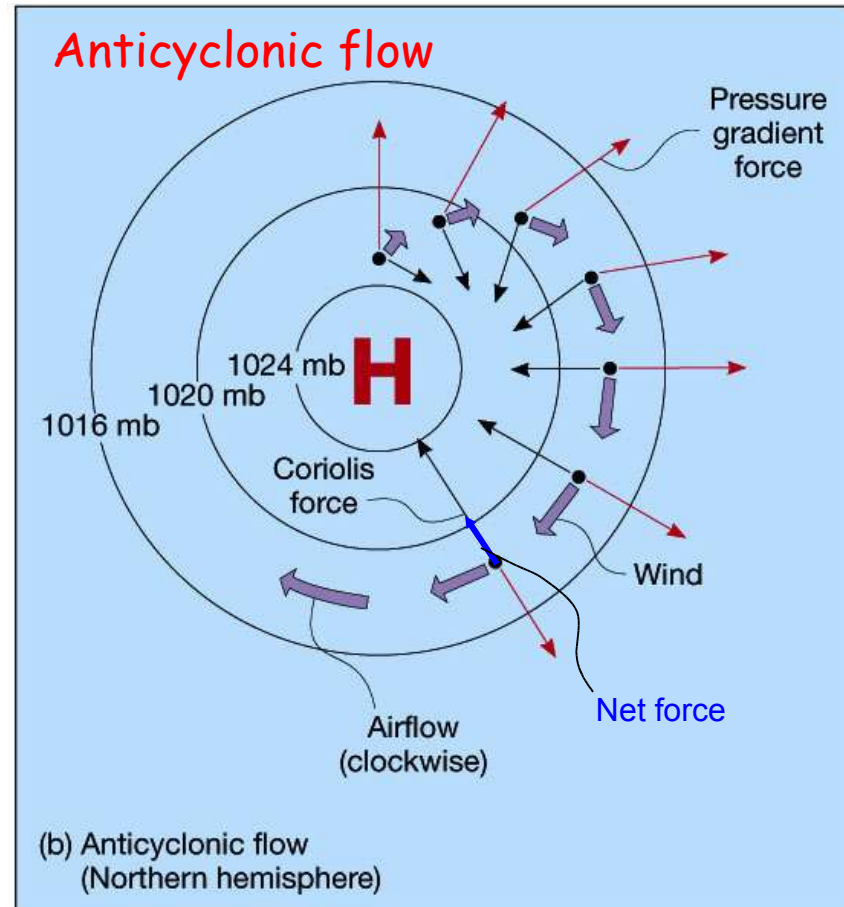
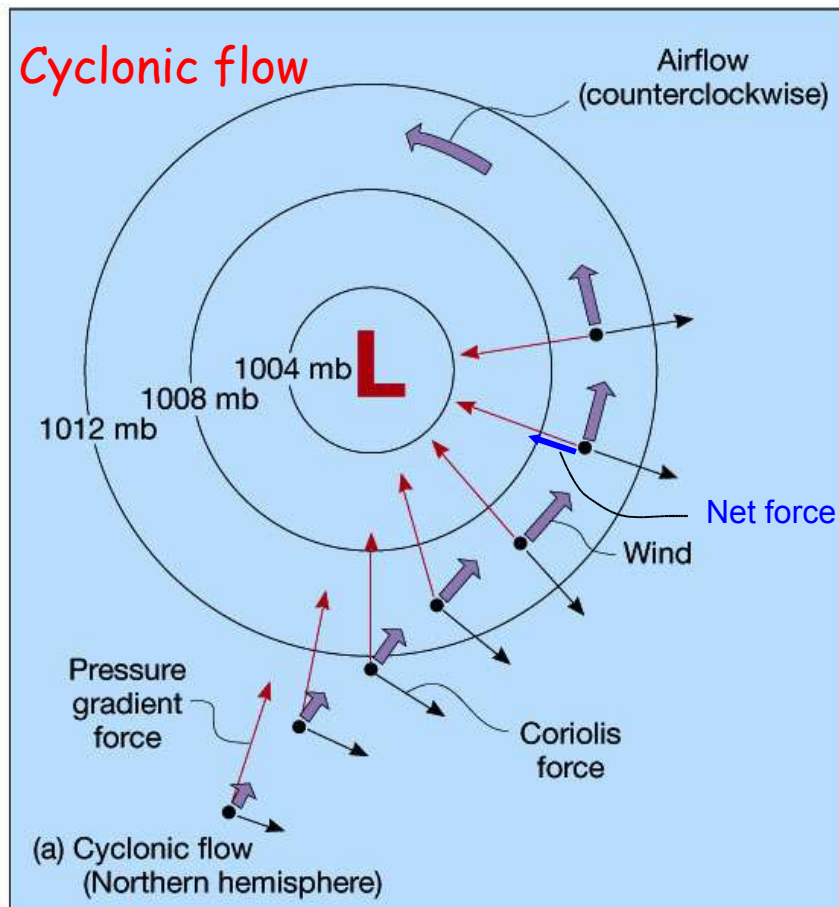
Gradient Winds (Northern Hemisphere)

$$F_{\text{pressure}} > F_{\text{coriolis}}$$

$$F_{\text{pressure}} < F_{\text{coriolis}}$$

The net force acts as a centripetal force.

$$F_{\text{net}} = F_{\text{pressure}} - F_{\text{coriolis}}$$

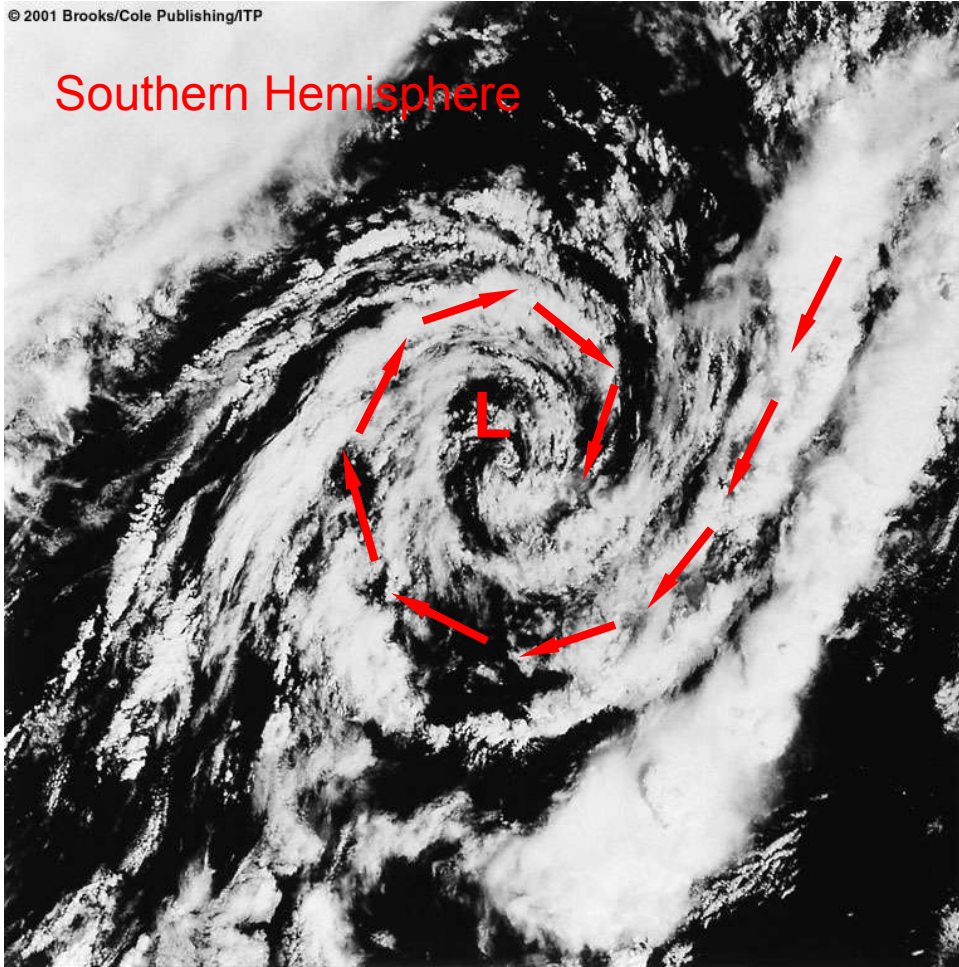


Cyclonic Flow (flow around a low pressure center)

- Clockwise in SH

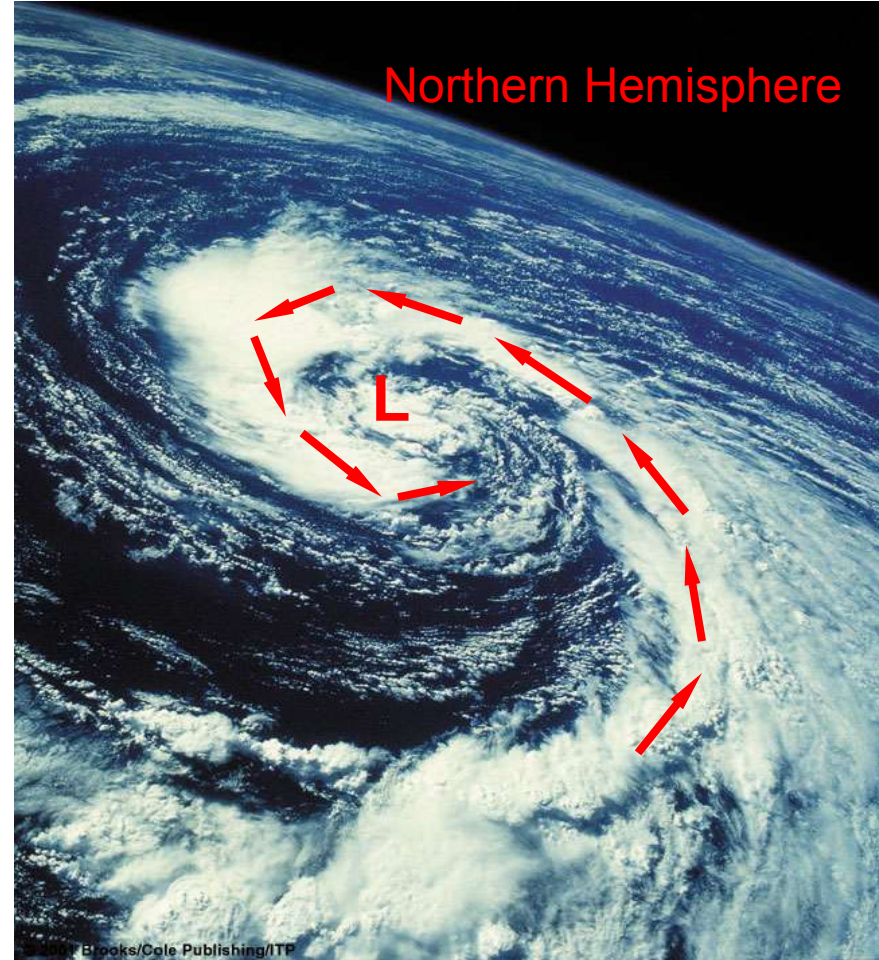
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Southern Hemisphere



- Counterclockwise in NH

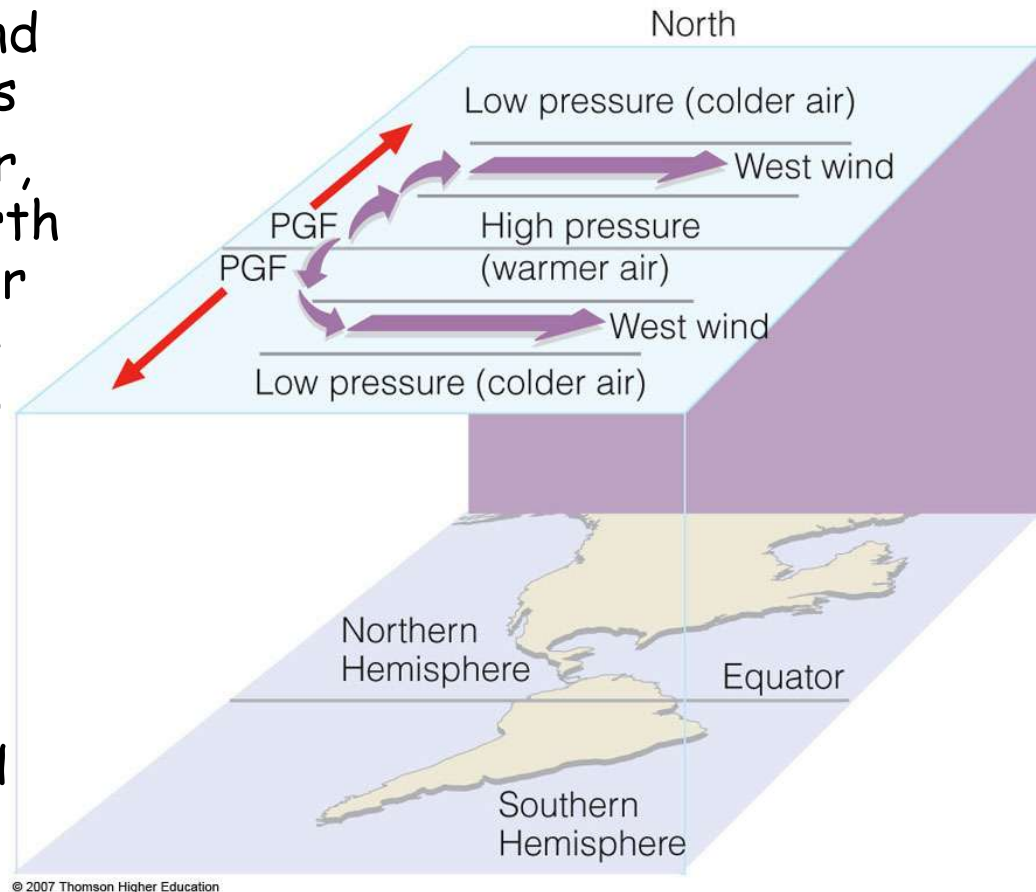
Northern Hemisphere



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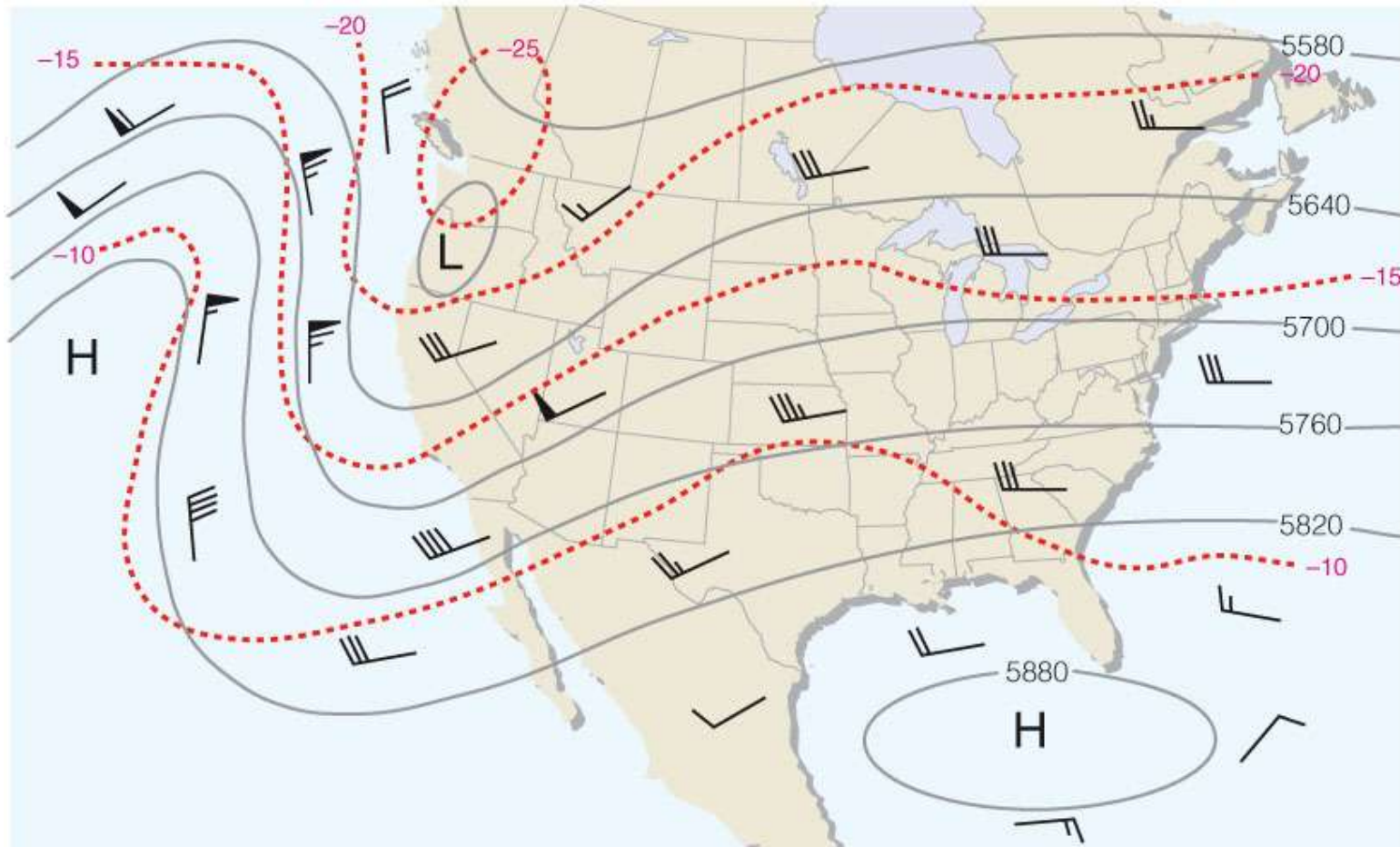
Winds Aloft in the Southern Hemisphere

- Warm air above the equator and cold air above the polar regions
- Higher pressure at the equator, lower pressure both to the north and to the south of the equator
- The **pressure gradient force** is towards the poles, sets the air in motion
- The **Coriolis force**
 - ◆ NH: to the right
 - ◆ SH: to the left
- The wind turns right in the NH and left in the SH, becomes parallel to the isobars
- **Westerly** winds in both the Northern and Southern Hemispheres.



Summary: prevailing winds at high altitudes

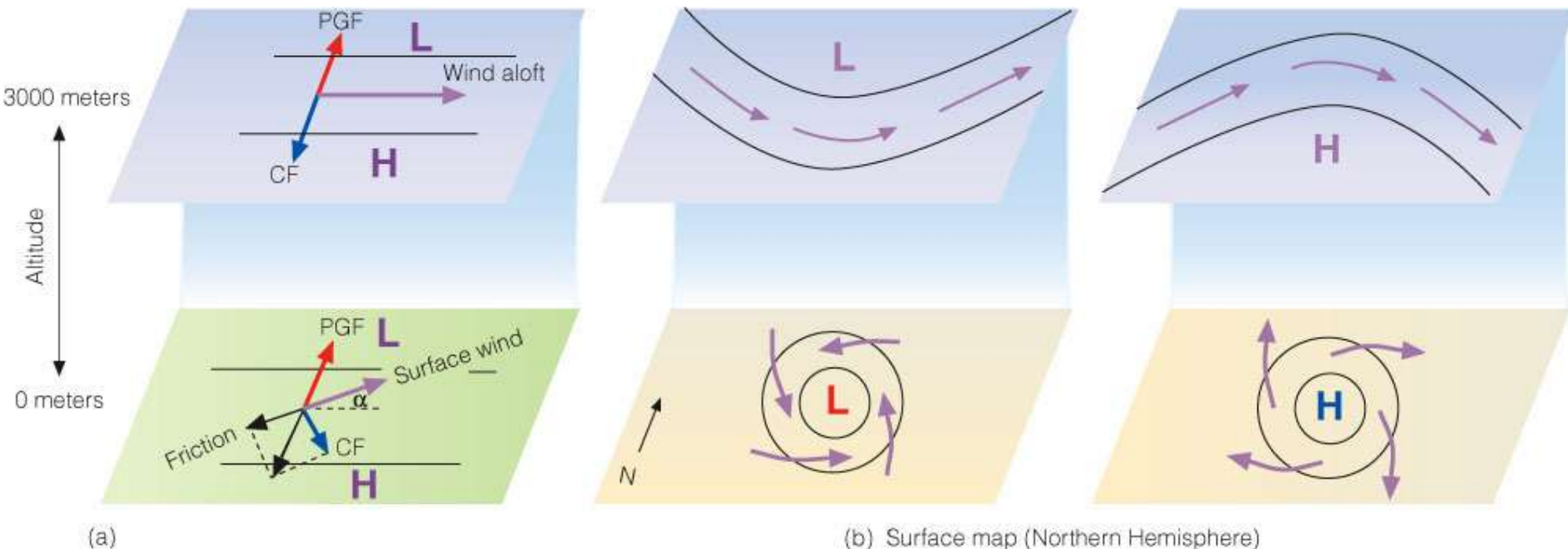
- Direction
 - ◆ Zonal: E-W
 - ◆ Meridional: N-S
- Balance of forces
 - ◆ Geostrophic: near straight isobars
 - ◆ Gradient: near curved isobars



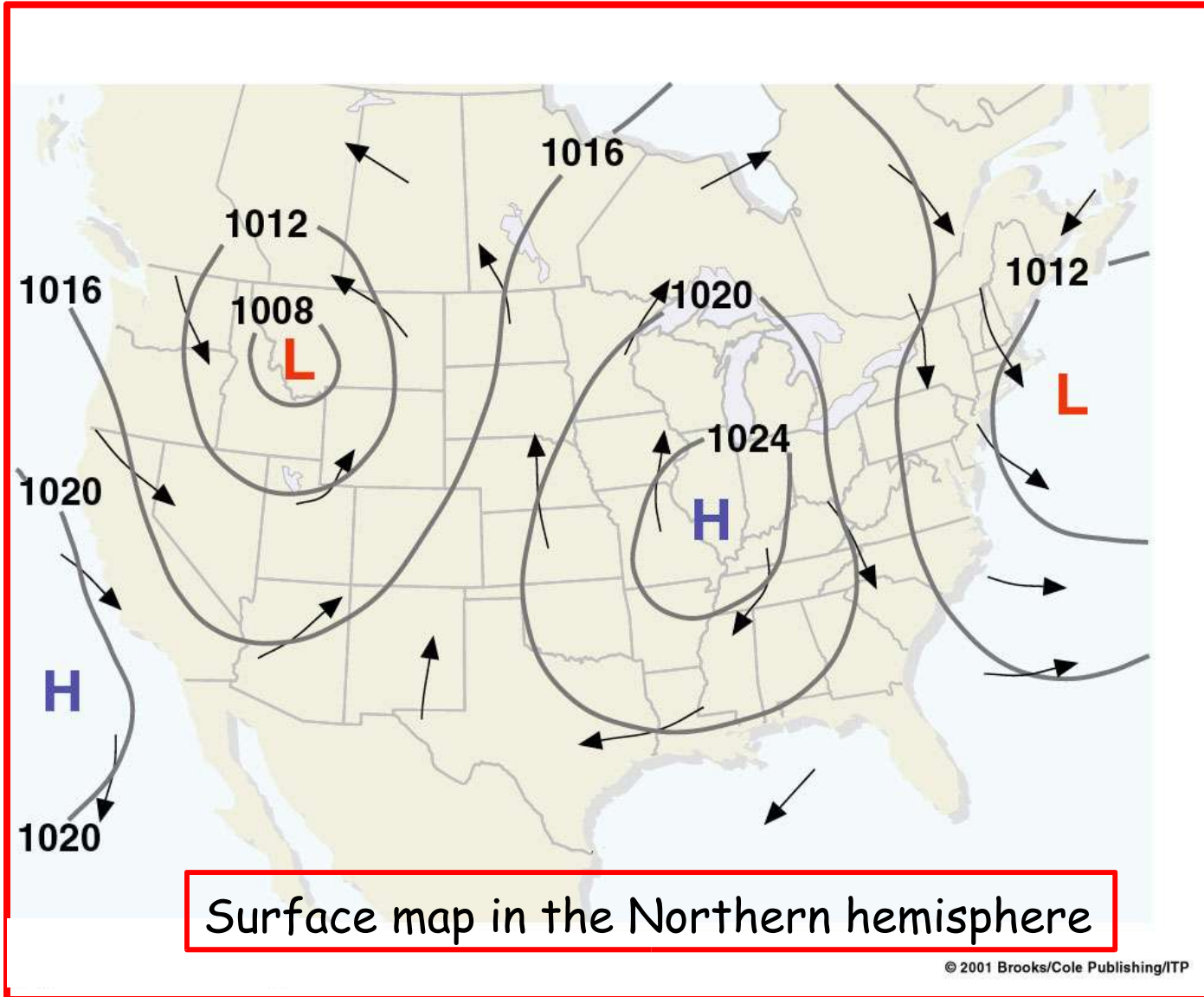
	Miles (statute) per hour	Knots
☉	Calm	Calm
—	1–2	1–2
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└└└	15–20	13–17
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└└└└└	26–31	23–27
└└└└└└	32–37	28–32
└└└└└└└	38–43	33–37
└└└└└└└└	44–49	38–42
└└└└└└└└└	50–54	43–47
└└└└└└└└└└	55–60	48–52
└└└└└└└└└└└	61–66	53–57
└└└└└└└└└└└└	67–71	58–62
└└└└└└└└└└└└└	72–77	63–67
└└└└└└└└└└└└└└	78–83	68–72
└└└└└└└└└└└└└└└	84–89	73–77
└└└└└└└└└└└└└└└└	119–123	103–107

Surface Winds-a balance of three forces

- In the boundary layer (~1km thick) **friction is important!**
- Friction is acting **opposite** the direction of the velocity -> friction **reduces** the wind speed -> the Coriolis force becomes **weaker** -> it **cannot balance** the pressure force.
- The wind starts to blow **across the isobars** towards the **low pressure**
- The angle between the direction of the wind and the isobars is on average 30 deg. It depends on the topography.

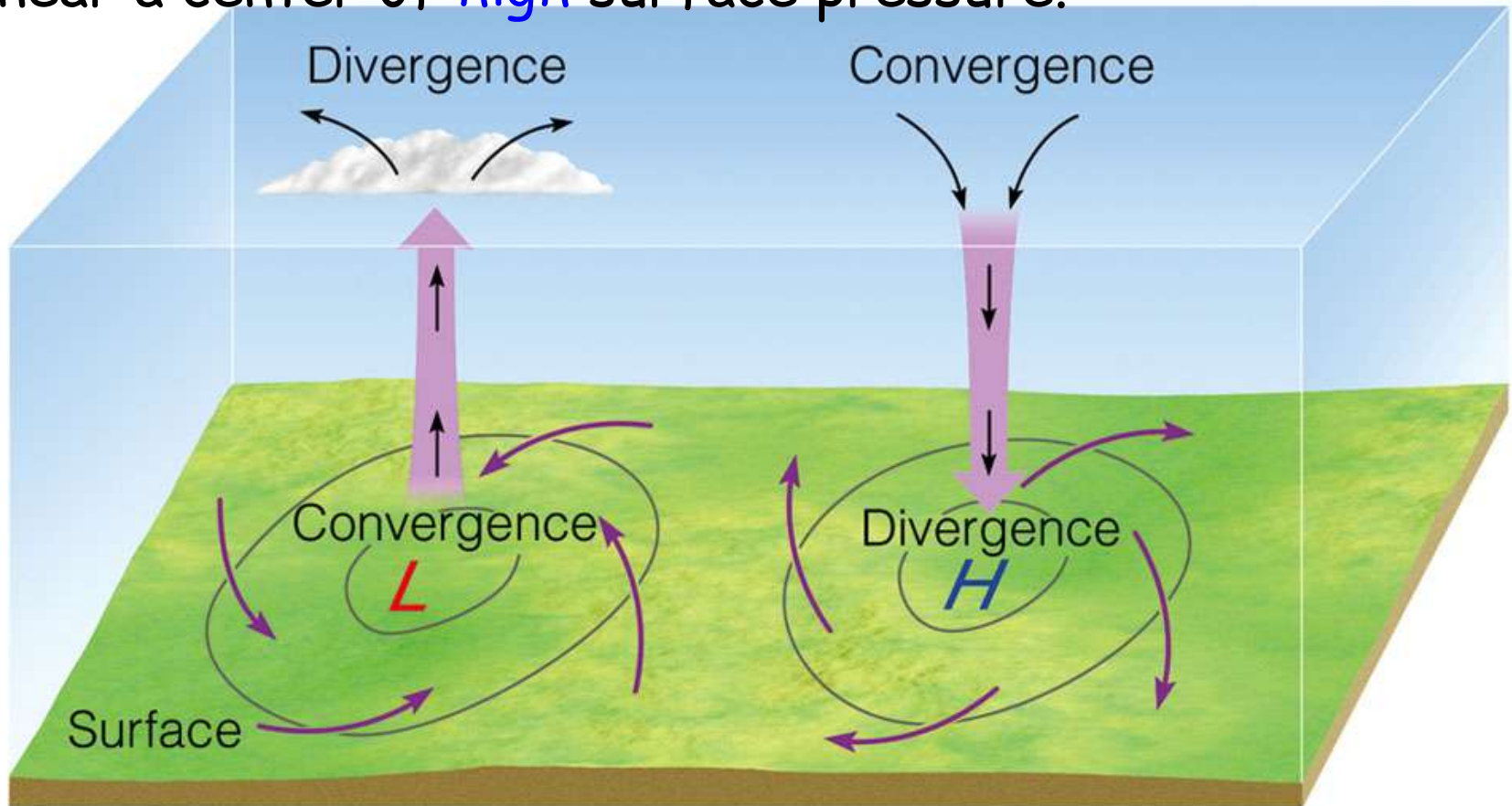


Is this a surface or a high-altitude map? Which hemisphere is this?



Vertical Air Motion: Convergences and Divergences

- Near a center of **low** surface pressure there is a **convergence** of air -> the air is forced to **rise** and then **diverge** at higher altitudes. The opposite takes place near a center of **high** surface pressure.



Hydrostatic Equilibrium

- On average **gravity** is balanced by the **pressure gradient force** -> **hydrostatic equilibrium**
- Small deviations from hydrostatic equilibrium result in small vertical winds (**a few cm/s**)

