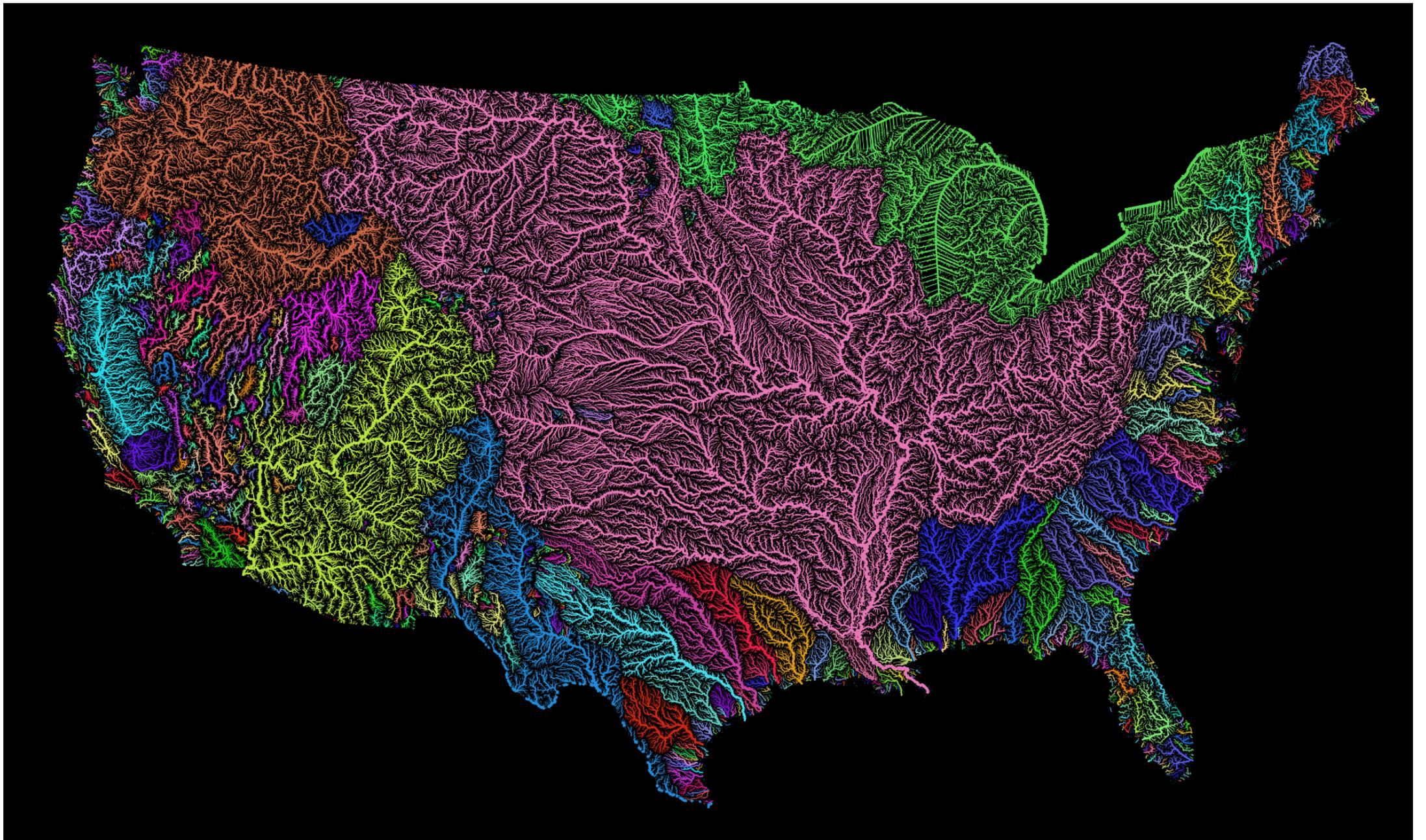


A Journal club presentation of Scaling and river networks: A Landau theory for erosion



Marc Lewkowitz



Source: <https://www.etsy.com/shop/GrasshopperGeography>

What is the goal of this model

🌊 How do rivers form?

≈ Water cycle and erosion

🌊 Can these fractal patterns and scaling laws be obtained from the laws of physics?

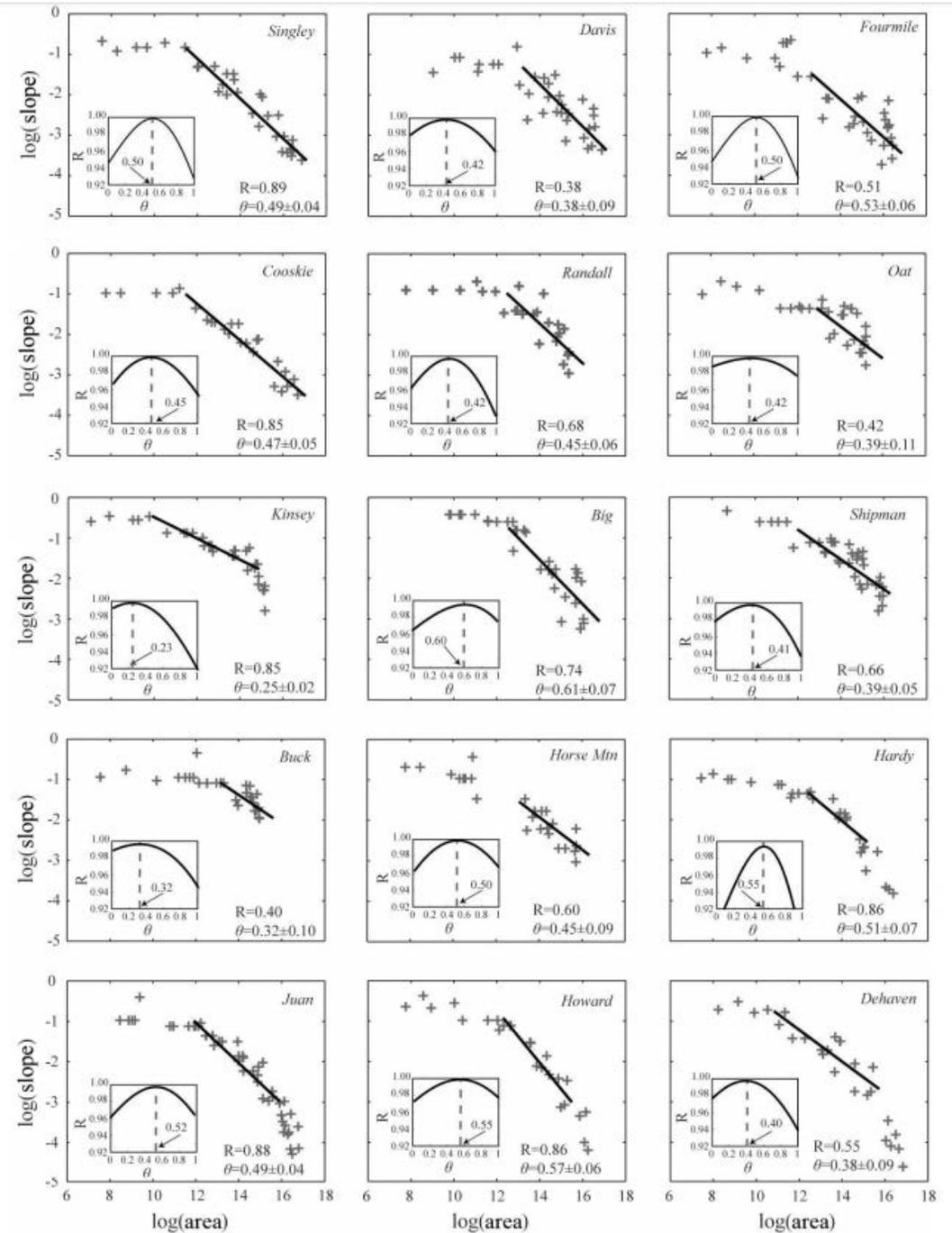
Universality in river networks

☞ Slope area law

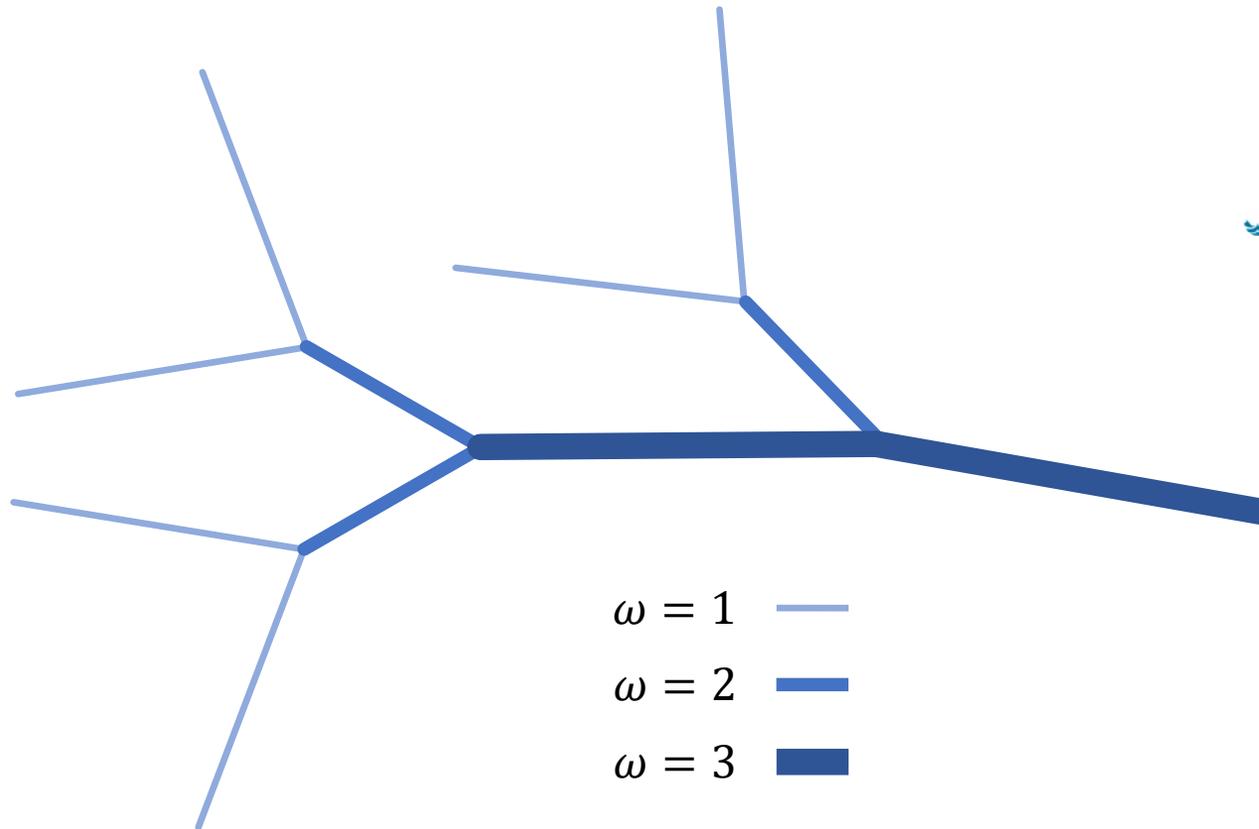
$$s = k_s Q^{-\theta}$$

$$\theta \approx 0.5$$

Slope area data from various rivers. $\theta=0.45$ on average
Taken by Wang, Y., Zhang, H., Zheng, D., Yu, J., Pang, J., and Ma, Y.



Horton's laws



☞ Horton's laws

$$R_b(\omega) = \frac{N_\omega}{N_{\omega+1}} \approx 4$$

$$R_L(\omega) = \frac{L_\omega}{L_{\omega+1}} \approx 2$$

☞ Fractal dimension

$$D = d_c \frac{\ln(R_b)}{\ln(R_L)}$$

$$d_c \approx 1.1 \text{ to } 1.2$$

The Model

☞ Assumptions

- ≈ Rainfall is uniform (no underground flow)
- ≈ Eroded material is not redeposited
- ≈ Water flows fast, soil flows slowly
- ≈ Initial landscape is without features

☞ Langevin equation

$$\dot{h} = F(h, h^2, \dots, \nabla h, \nabla^2 h, \|\nabla h\|^2, \dots) + \eta(\mathbf{r}, t)$$

Simplifying.

- ☞ The change in the heights shouldn't depend on the height directly

$$\dot{h} = F(\nabla h, \nabla^2 h, \|\nabla h\|^2, \dots) + \eta(\mathbf{r}, t)$$

$$F = A + \mathbf{B} \cdot \nabla h + C * \|\nabla h\|^2 + D * \nabla^2 h + \dots$$

- ☞ A Represents a uniform motion of the entire system ($A = 0$)

- ☞ \mathbf{B} Represents a preferred direction of soil flow ($\mathbf{B} = 0$)

$$\dot{h} = C * \|\nabla h\|^2 + D * \nabla^2 h + \eta(\mathbf{r}, t)$$

≈ Kardar-Parisi-Zhang (KPZ) Equation

- ☞ C Is the erosion coefficient

- ☞ D Is a smoothing term

Just add water

- ☁ Uniform rainfall

$$\nabla \cdot \mathbf{q} = R$$

- ☁ Water flows down hill (Dynamical system approximation)

$$\mathbf{q} = \nabla h * \sigma$$

- ☁ No water means no erosion

$$C = -c * \|\mathbf{q}\|$$

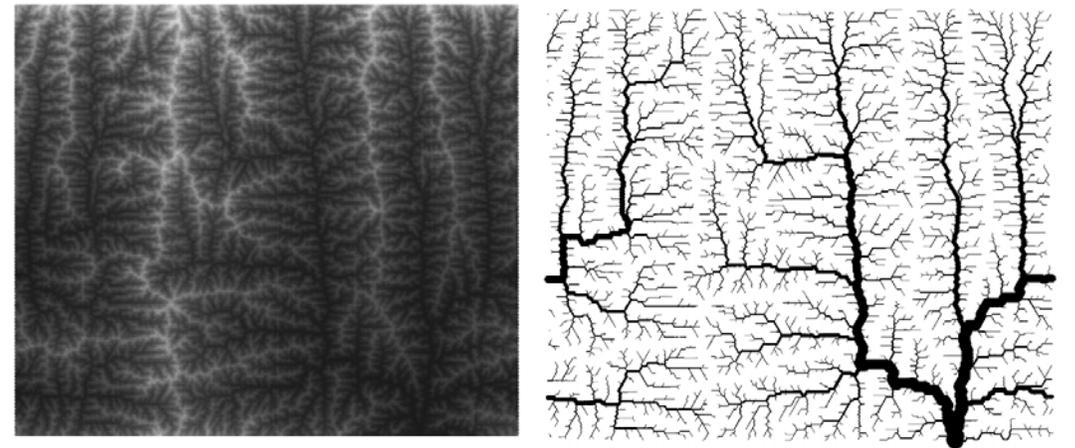
$$\dot{h} = -c * \|\mathbf{q}\| \|\nabla h\|^2 + D * \nabla^2 h + \eta(\mathbf{r}, t)$$

- ☁ A final state implies uniform erosion.

$$k = -c * \|\mathbf{q}\| \|\nabla h\|^2 \Rightarrow \|\nabla h\| \propto \|\mathbf{q}\|^{-1/2}$$

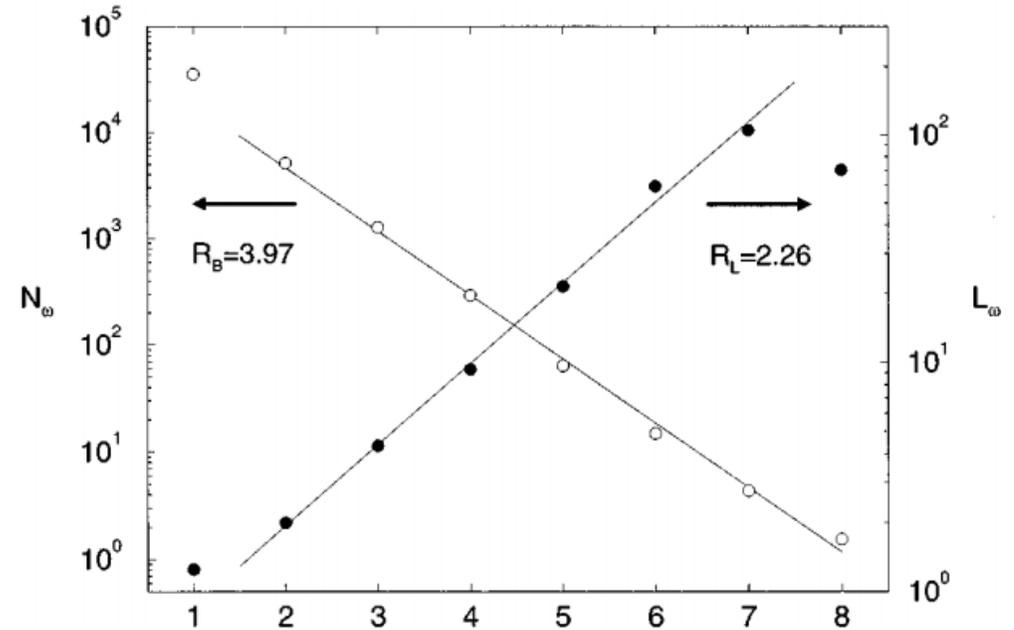
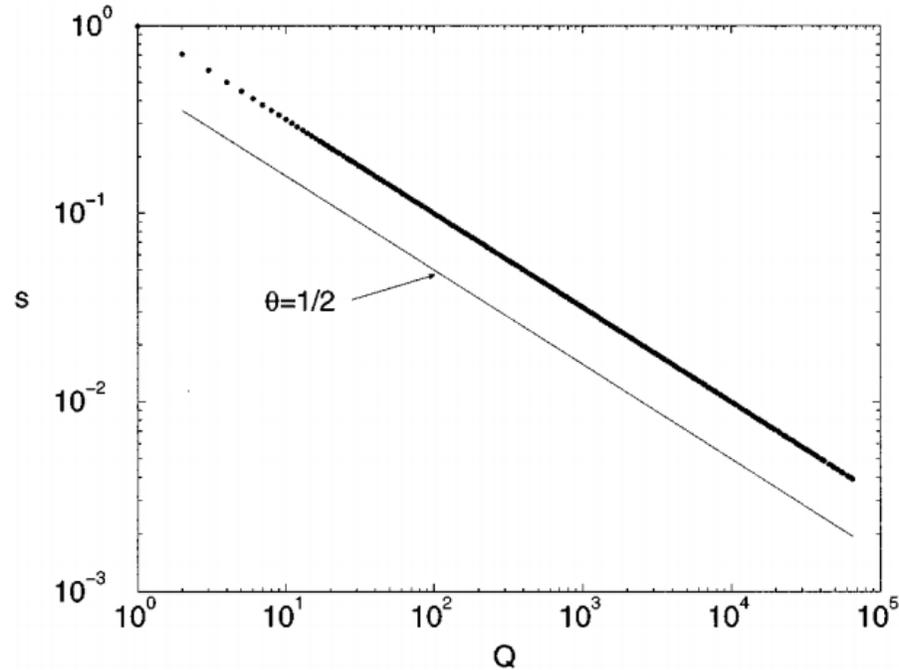
Simulation

- 🌊 256x256 Triangular lattice
- 🌊 Water flows along edges between lattice points
- 🌊 Initial landscape is a downward sloping hill with some noise
 - ≈ No local minima in the initial lattice
- 🌊 Updates occur according to
$$\Delta h = -\|\nabla h\|^2 \|q\| \Delta t$$



Sample data from a single simulation, the left is a height map and right is the water flow
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The results



Results from 20 independent simulations, The right suggests that $\theta = \frac{1}{2}$ is an attractive solution. The left shows that the simulation results are similar to that of real rivers With respect to Horton's laws.

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Citations

- ☩ Somfai, E. & Sander, L. M. (1997). Scaling and river networks: A Landau theory for erosion. *Phys. Rev. E*, 56, R5-R8.
- ☩ Wang, Y., Zhang, H., Zheng, D., Yu, J., Pang, J., and Ma, Y.: Coupling slope–area analysis, integral approach and statistic tests to steady-state bedrock river profile analysis, *Earth Surf. Dynam.*, 5, 145-160, <https://doi.org/10.5194/esurf-5-145-2017>, 2017.